

Physicochemical Characterization of Pond Waters in the Kanem Region, Republic of Chad

Jean-Paul Vicat¹, Jean-Claude Doumnang Mbaigane¹, Nadjilem Dingamtar Ndjadodi¹

Laboratory of Geology, Geomorphology, and Remote Sensing, Faculty of Exact and Applied Sciences, University of N'Djamena, P.O. box 1027, N'Djamena, Republic of Chad

Email address: jpfv@hotmail.fr

Abstract— The physicochemical properties of ten water samples from ponds in the Kanem region and a sample from the southern part of Lake Chad were analyzed. Surface water samples were collected in October 2007. The following parameters were determined: temperature, pH, electrical conductivity, the major ions (Ion *Chromatography), SiO₂ and HCO₃⁻ (UV-visible Spectrophotometer)* and trace elements (Li, Rb, Cs, Be, Sr, Ba, Al, V, Cr, Mn, Co, Ni, Cu, Zn, Mo, Ag, Cd, Sn, Pb, Bi, As, Sb, U) (ICP-MS). Based on pH the waters are neutral or alkaline. The waters exhibited low, medium, and high salt enrichments based on conductivity. Compared to Lake Chad, the concentrations of major ions vary significantly. Salinity, defined by Total Dissolved Solids, ranged from fresh, brackish, and saline to brine waters. Water chemistry is predominantly influenced by the evaporation-crystallization process. In the Piper and Chadha diagrams, the pond waters are mainly of the Na-HCO3 type. Trace element concentrations are either enriched or depleted compared to Lake Chad and resemble those of soda or saline lakes. Occasionally elevated concentrations of As (0.20-0.70 mg/L) could impact the quality of harvested spirulina from the Kanem ponds.

Keywords—Chad, Kanem, Pond, Water, Major and Trace Elements, Arsenic, Spirulina.

I. INTRODUCTION

To the northeast of Lake Chad, the Kanem region is occupied by a fossil erg composed of dunes-oriented NE-SW. The climate is sub-desertic with annual rainfall of less than 300 mm. Temperatures are high, with an average of around 28°C annually. The interdunal depressions are often occupied by permanent or temporary ponds locally referred to as "ouadi." The term "ouadi," derived from "oued," is improper as it implies temporary surface drainage, whereas the ouadis of Kanem have neither watershed nor outlet. The ponds are solely supplied by sources from the shallow groundwater table and rainfall. In Chad, the term "lake" is used for large permanent ponds. Similar to the ponds, lakes are shallow, with depths of about a meter. This renders the term lake incorrect as per Touchart's (2000) terminology, but it has been commonly used since early studies (Iltis, 1969), in agreement with Dussart's (1966) nomenclature. The sub-desert climate leads to the precipitation of evaporitic minerals in the ouadis, primarily gaylussite and trona (Maglione, 1968), as well as Mg-calcite, natron, nahcolite, thermonatrite, halite, northupite, gypsum, mirabilite, thenardite, bloedite, magadiite, kenvaite, kanemite, mordenite (Maglione, 1971, 1976; Maglione and Karn, 1975). Under favorable water salinity conditions (Iltis, 1968, 1969), spirulina develops in the ponds of the ouadis. The ouadis are centers of intense activity for subsistence farming and sometimes for trona exploitation and spirulina harvesting. In the 1950s–1970s, the Kanem region was the subject of numerous scientific studies, but the chemical analyses of ouadi waters conducted during these studies (Iltis, 1971a, 1971b; Maglione, 1969; 1976) only focused on major ions. In this paper, we present and discuss new geochemical results regarding the analysis of major ions and trace elements in the waters of ten ouadis that have not been studied previously.

II. GEOLOGICAL SETTING

The Kanem region is occupied by a Quaternary fossil erg. The aeolian sand formations were deposited during the Pleistocene Last Glacial Maximum approximately 200,000 years ago (Servant, 1983; Maley, 2010; Sylvestre, 2014). During the Holocene, between ca. 12,000 and ca. 5,000 years BP (Amaral et al., 2013; Sylvestre et al., 2019), the climate becomes wetter. Lake Chad then extended to a coastline at 325 meters and formed the Mega Lake Chad (Fig. 1a), covering an area of 350,000 km2 (Schuster et al., 2005; Leblanc et al., 2006a, b). During this transgression, the Mega Lake Chad partially covered the aeolian sand formations, and the lacustrine Labdé series (Servant and Servant, 1970) was deposited in interdunal depressions. This series consists of clay and silts (Dupont, 1967). After around ca. 5,000 years BP, the climate becomes more arid, resulting in a reduction in the surface area of Lake Chad (Sylvestre et al., 2019), which is today estimated at around 25,000 km² (CBLT 2023). The Labdé series now outcrops, forming the bottom of the ouadi ponds.

The existence of the ponds depends on rainfall and groundwater levels. The persistence of the ponds is possible when water inputs from the shallow groundwater compensate for the deficit between precipitation and evaporation. The piezometric surface of the groundwater is marked by a depression (Fig. 1b) along the northeast border of Lake Chad (Schneider, 1967 a, b). The ouadis located southwest of this depression are fed by the Lake Chad aquifer; their water levels depend on Lake Chad's water level, which is influenced by rainfall in its watershed. The ouadis located in the northeast of the piezometric depression are fed by the Chitati aquifer, with the underground water divide located about fifty kilometers northeast of Lake Chad. Their water level depends solely on the local rainfall regime that feeds the Chitati aquifer in the studied region.



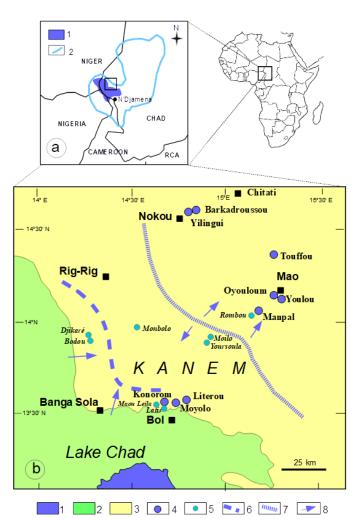


Fig.1. Schematic map of the study area. (a) Mega Lake Chad. 1: Current Lake Chad. 2: Holocene Mega Lake Chad. (b) Location map of the ouadis. 1: Lake Chad, permanent open water. 2: Lake Chad, polders. 3: Aeolian sand formations and undifferentiated Quaternary deposits. 4: Studied ouadis. 5: Ouadis mentioned in the text. 6: Hydrogeological depression. 7: Groundwater divide. 8: Groundwater flow direction.

Over the past 5,000 years, the transition to a more arid climate in the Lake Chad basin has not been linear and has shown alternating wet and dry periods (Maley, 2010). The current period is characterized by significant irregularities in the annual precipitation regime (Maglione, 1969), leading to cycles of drying and refilling of the ouadi ponds. As such, the ouadis have a complex history marked by several hydrogeological cycles.

Among the ouadis that have been the subject of previous studies (Iltis, 1971a, 1971b; Maglione, 1969; 1976), only the Monbolo and Yoursoula lakes are still in water. The Bodou, Rombou, Djikaré, and Moilo lakes and the Latir and Maou-Leila ponds are currently dried up.

III. SAMPLES COLLECTION AND ANALYSIS

Water samples for analysis were collected in October 2007. The samples were taken from the Oyouloum, Mampal, Moyolo, Rombou, Literou, Youlou, Konorom, Barkadroussou, Touffou, and Yilingui ouadis. The locations of the ouadis are shown in Figure 1b. For comparison, a sample was also taken from the southern part of Lake Chad at Lat 12.996198° N and Long 14.576358° E.

The physicochemical parameters, pH, and temperature were measured in situ using the pH 323 / Set B mixed electrode, and conductivity was measured using the Lf 318 / Set electrode. Several water samples were collected at each site for analysis, and the following sample preparation protocols were conducted in the field. The waters were initially filtered using Millipore mixed cellulose 0.45 μ m filters through a pressure vessel and a Millipore Swinnex filter holder. The samples were stored in precleaned high-density polyethylene bottles. Samples for trace element analysis were then filtered at 0.22 μ m. Water samples for cations and trace elements analysis were acidified with ultra-pure nitric acid.

Major ions analysis was performed at the Laboratory of Hydrogeology at the University of Avignon, France. Major ions were analyzed using Ion Chromatography. The overall detection limit for ions was 0.04 mg/L. SiO_2 and HCO_3^- were analyzed using a UV-visible Spectrophotometer. Trace element analyses were conducted at the Hydro-Sciences Laboratory, University of Montpellier II (France) using Inductively Coupled Plasma-Mass Spectrometry with a relative uncertainty of less than 5%. Quality control was performed using international water standards

IV. RESULTS AND DISCUSSION

The physical and chemical characteristics of the ouadi waters are presented in Table 1.

4.1 Physical Parameters

The temperature of the ouadi waters is variable. The Barkadroussou and Konorom ouadis have temperatures of 24.30°C and 24.40°C, respectively, similar to Lake Chad measured at 23.40°C. The other ouadis have higher temperatures ranging from 27.40°C to 33.30°C.

The pH values of the ouadi water samples range from 7.45 to 10.17. According to the WHO classification (2017), the waters of Oyouloum and Manpal ouadis with pH values between 6.5 and 8.5 are neutral. The waters of the other ouadis, with pH values above 8.5, are alkaline.

The electrical conductivity (EC) ranged from 227 to 14,800 μ S/cm. EC is a measure of the ionic concentration (i.e., salinity) of the water. EC is classified according to Sarath Prasanth et al. (2012) into Type I (EC < 1,500 μ S/cm), Type II (1,500 < EC < 3,000 μ S/cm), and Type III (EC > 3,000 μ S/cm). According to this classification, the Oyouloum and Manpal ouadis fall under Type I (low enrichment of salts). The Moyolo and Rombou ouadis fall under Type II (medium enrichment of salts), and the other ouadis fall under Type III (high enrichment of salts).

4.2 Major ions

The Oyouloum ouadi has a low Na^+ concentration (16.67 mg/L), similar to that of Lake Chad (17.50 mg/L). The other ouadis have highly variable Na^+ concentrations. The Manpal,



	1		nysicoenen	neur purum	eters of out	di waters a	ila a sumple	i i o in the B		rt of Lake C	inud.	
		Oyouloum	Mampal	Moyolo	Rombou	Literou	Youlou	Konorom	Barkadroussou	Touffou	Yilingui	Lake Chad
T°C		30.70	28.60	27.40	32.70	26.60	33.30	24.40	24.30	29.00	29.90	23.40
pН		7.45	7.73	9.36	9.30	9.61	9.39	9.97	10.09	9.87	10.17	7.60
c 25°C		227	460	2,170	2,700	5,870	21,500	57,000	72,400	75,600	94,800	335
Na+	mg/L	16.97	55.92	441.60	565.50	1,252	5,365	21,840	25,630	27,950	35,550	17.50
K+	mg/L	5.13	12.83	124.40	146.30	334.80	1,276	1,292	6,053	6,273	11,050	8.60
Ca++	mg/L	24.07	32.49	11.33	3.00	2.80	15.80	8.75	7.25	10.25	8.25	16.20
Mg++	mg/L	1.79	1.82	3.79	2.02	8.10	15.60	57.00	55.50	98.75	44.50	7.50
HCO3-	mg/L	105.75	184.18	968.93	1,130	2,120	7,596	16,604	30,715	40,715	27,534	121.40
CO3	mg/L	0.12	0.41	91.02	92.45	354.12	764.50	6,354	15,494	12,376	16,699	1.80
NO3-	mg/L	< 0.03	0.52	< 0.03	< 0.03	< 0.03	< 0.03	< 0.03	< 0.03	28.36	< 0.03	-
NO2-	mg/L	0.55	9.25	< 0.02	< 0.02	< 0.02	< 0.02	< 0.02	< 0.02	< 0.02	< 0.02	-
Cl-	mg/L	3.91	7.50	112.57	117.96	470.90	1,789	874.88	6,841	6,027	16,653	4.30
SO4	mg/L	19.43	52.94	13.51	135.14	146.17	2,869	19,578	784	2,557	4,414	6.20
PO4	mg/L	< 0.03	0.66	11.12	12.42	37.31	204.52	254.56	674.58	572.14	1,247	4.27
Br-	mg/L	< 0.03	0.18	1.53	2.31	9.38	31.53	45.55	100.92	86.68	296.13	-
F-	mg/L	0.07	0.19	0.48	0.33	1.20	1.20	2.00	130.87	533.43	7.84	0.90
SiO2	mg/L	58.57	87.21	115.04	79.91	101.41	198.59	209.95	248.00	315.58	333.91	48.54
ionic B	0	-0.01	-0.01	0.01	0.01	0.00	0.01	0.03	0.01	0.01	0.07	-0.01
Li	μg/L	1.26	0.92	3.08	1.07	1.83	1.81	1.89	16.54	27.75	20.05	0.30
Rb	μg/L	1.83	1.65	16.58	51.20	238.40	383.50	148.10	3,691	4,437	7,642	6.56
Cs	μg/L	0.003	0.01	0.01	0.02	0.06	0.07	0.02	0.95	1.25	1.37	0.02
Be	μg/L	< 0.003	< 0.003	< 0.003	< 0.003	< 0.003	0.03	0.03	0.05	0.05	0.05	< 0.003
Sr	μg/L	89.74	120.30	256.40	44.10	137.00	97.97	395.10	181.90	141.70	144.10	91.23
Ba	μg/L	32.76	29.82	176.80	22.75	255.80	63.22	157.70	196.70	226.90	228.70	70.38
Al	μg/L	< 0.129	< 0.129	0.13	<0.129	0.13	5.68	2.33	8.16	15.04	8.48	47.87
V	μg/L	1.33	0.42	7.83	30.85	26.25	95.78	228.70	85.40	21.05	245.90	3.88
Cr	μg/L	< 0.012	< 0.012	1.69	1.11	9.67	41.92	33.40	65.65	92.12	100.80	1.00
Mn	μg/L	< 0.003	0.01	0.63	0.58	2.31	9.30	6.07	23.77	64.65	9.91	3.43
Co	μg/L	0.09	0.07	0.14	0.08	0.41	2.21	1.22	1.53	9.98	1.03	0.28
Ni	μg/L	0.08	0.26	0.15	0.45	5.73	51.66	78.61	3.20	10.02	18.00	1.89
Cu	μg/L	0.08	0.81	0.48	3.26	4.09	7.93	25.64	29.78	24.25	60.51	2.18
Zn	μg/L	<0.159	< 0.159	0.97	<0.159	5.27	7.18	14.99	5.57	9.94	9.77	16.41
Mo	μg/L	0.75	2.11	2.77	3.64	16.26	206.00	483.90	378.10	82.91	691.90	0.43
Ag	μg/L	< 0.009	< 0.009	< 0.009	< 0.009	< 0.009	< 0.009	< 0.009	< 0.009	0.04	< 0.009	< 0.009
Cd	μg/L	< 0.003	< 0.003	0.003	< 0.003	0.84	0.25	1.65	1.64	0.64	0.49	0.13
Sn	μg/L	0.68	0.46	0.29	0.80	0.94	3.63	7.80	11.17	47.38	2.26	5.00
Pb	μg/L	< 0.003	< 0.003	0.003	< 0.003	0.38	< 0.003	0.96	1.10	6.13	< 0.003	0.19
Bi	μg/L μg/L	25.59	35.29	100.40	141.80	774.60	2,939	3,018	7,054	5,905	11,400	5.69
As	μg/L	0.49	0.61	1.51	11.49	18.54	201.30	697.70	377.90	322.10	640.10	0.04
Sb	μg/L μg/L	0.04	0.002	0.04	0.12	0.24	1.98	8.42	2.22	1.57	3.55	0.06
U	μg/L	0.04	0.03	0.09	1.02	3.49	6.89	201.40	28.26	15.30	32.18	0.27
TDS	μg/L mg/L	237	446	1,896	2,288	4,846	20,129	67,127	85,222	97,556	113,858	190
105	mg/L	231	440	1,070	2,200	4,040	20,129	07,127	05,222	21,550	115,656	190

TABLE 1. Physicochemical parameters of ouadi waters and a sample from the Southern part of Lake Chad.

Moyolo, Rombou, Literou, and Youlou ouadis have concentrations (55.92–5,365 mg/L) lower than that of seawater estimated at 10,784 mg/L (UNESCO, 1994), while the other ouadis have significantly higher concentrations (21,840–35,365 mg/L) than that of seawater.

The Oyouloum and Manpal ouadis have K^+ concentrations of 5.13 mg/L and 12.83 mg/L, respectively, similar to that of Lake Chad (8.6 mg/L). The other ouadis have higher and more variable K^+ concentrations (124.4–11,050 mg/L).

The Ca^{2+} concentrations in the Oyouloum and Manpal ouadis are 24.07 and 32.49 mg/L, respectively, higher than that

of Lake Chad (16.20 mg/L). The other ouadis have lower Ca^{2+} concentrations (2.80–15.80 mg/L).

The Oyouloum, Manpal, Moyolo, and Literou ouadis have Mg^{2+} concentrations (1.79–2.02 mg/L) lower than that of Lake Chad (7.50 mg/L). The other ouadis have higher Mg^{2+} concentrations (8.10–98.75 mg/L).

The concentrations of HCO₃⁻ (105.75–40,715 mg/L) and CO₃²⁻ (0.12–5,494 mg/L) are highly variable. The HCO₃⁻ concentrations in the Oyouloum and Manpal ouadis, 105.75 and 184.18 mg/L, respectively, are similar to that of Lake Chad (121.4 mg/L). The other ouadis have significantly higher HCO₃⁻ concentrations (968.93–40,715 mg/L). The CO₃⁻²



concentrations in the Oyouloum and Manpal ouadis, 0.12 and 0.41 mg/L, respectively, are lower than that of Lake Chad (1.80 mg/L). The other ouadis have significantly higher CO_3^{2-} concentrations (91.02–16,699 mg/L).

 NO_3^- concentrations are below the detection limit (0.03 mg/L), except for the Manpal ouadi (0.52 mg/L) and the Touffou ouadi (28.36 mg/L). Nitrate concentrations in the Touffou ouadi are high compared to the other ouadis, where concentrations are similar to most saline lakes (< 1 mg/L) (Jirsa et al., 2013). Nitrate (NO_3^-) and nitrite (NO_2^-) concentrations result from the nitrification of ammonium ions (NH_4^+) derived from incomplete degradation of organic matter in the water. Ammonium ions are oxidized to nitrites and then to nitrates. The elevated NO_3^- values in the Touffou ouadi may be attributed to the high density of spirulina during sample collection (Doumnang Mbaigane, 2008), as spirulina has high rates of decomposition and subsequent nitrification (Hammer, 1986).

NO₂⁻ concentrations in the ouadis are below the detection limit (0.02 mg/L), except for the Oyouloum and Manpal ouadis, which have concentrations of 0.55 mg/L and 9.25 mg/L, respectively. Nitrites are unstable and naturally occur in low quantities in waters (Appelo and Postma 1996; OEHHA, 1997). The presence of nitrites may be attributed to a significant influx of nitrogen (Forrest et al., 2006; Dubrovsky et al., 2010). We attribute the presence of nitrites in the Oyouloum and Manpal ouadis to recent pollution from the excreta of livestock drinking in these ouadis.

The Cl⁻ concentrations in the Oyouloum and Manpal ouadis are low, 3.91 and 7.5 mg/L, respectively, similar to that of Lake Chad (4.30 mg/L). The other ouadis have significantly higher Cl⁻ concentrations (112.57-16,653 mg/L).

 SO_4^{2-} concentrations are highly variable (13.51-1,958 mg/L) and higher than that of Lake Chad (6.2 mg/L). The presence of SO_4^{2-} in waters is due to gypsum deposits in the ouadis.

The Oyouloum and Manpal ouadis have very low PO_4^{2-} concentrations, <0.03 and 0.66 mg/L, respectively, lower than that of Lake Chad (4.27 mg/L). The other ouadis have highly variable PO_4^{2-} concentrations (11.12-1,247 mg/L), higher than that of Lake Chad. The PO_4^{2-} concentrations are attributed to the use of phosphate fertilizers for agricultural activities in the ouadis.

The Oyouloum and Manpal ouadis have low Br⁻ concentrations, <0.03 and 0.18 mg/L, respectively. The other ouadis have higher concentrations (2.31-296.13 mg/L).

The Barkadroussou and Touffou ouadis have high F^- concentrations, 130.87 and 533.43 mg/L, respectively, much higher than that of Lake Chad (0.90 mg/L). The other ouadis have lower (0.07-0.33 mg/L), similar (1.2-2 mg/L), or higher (7.84 mg/L) F^- concentrations than that of Lake Chad.

 SiO_2 concentrations range from 58.57 to 333.91 mg/L and are higher than that of Lake Chad (48.54 mg/L).

4.3 Salinity and Total Alkalinity

There are numerous descriptive terms in the literature for describing water salinity, and there are several different classification systems. Typically, Total Dissolved Solids (TDS) is used to define salinity. According to TDS (Kharaka and Hanor, 2003), waters are classified as freshwater (TDS < 1 g/L), brackish (1 < TDS < 10 g/L), saline (10 < TDS < 35 g/L), and brine (TDS > 35 g/L). The TDS values of the ouadi waters (Table 1), calculated by adding the total ion concentration of the water, are highly variable (205 to 11,386 mg/L). According to the above classification, the Oyouloum and Manpal ouadis, similar to Lake Chad (TDS = 167 mg/L), are classified as freshwater. The Moyolo, Rombou, and Literou ouadis are Brackish, the Youlou ouadi is Saline, and the Konorom, Barkadroussou, Touffou, and Yilingui ouadis are Brines.

The ouadi waters exhibit an excess of total alkalinity (TA = $HCO_3^- + 2(CO_3^{2-})$ i.e., the sum of the charges of the bicarbonate ion plus carbonate ion) over the charges of the alkaline earth ions magnesium and calcium: (HCO_3^-) + 2(CO_3^{2-}) > 2(Mg^{2+}) + 2(Ca^{2+}). According to Kempe and Kazmierczak (2011), this characterizes "soda lakes" where evaporation of the waters will lead to an increase in pH, and eventually, sodium carbonates will precipitate.

4.4 Mechanisms Controlling Water Chemistry

The Gibbs (1970) diagram is an effective tool for identifying the major mechanisms controlling natural water chemistry, including precipitation, rock-water interaction, and evaporation.

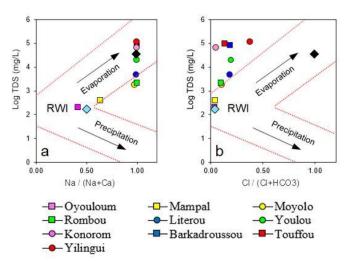


Fig. 2. Gibbs (1970) plots, a and b, showing the dominant factors controlling the ouadi water chemistry. RWI: Rock-Water interaction. Blue diamond: Lake Chad. Black diamond: Sea Water (UNESCO, 1994).

In the Gibbs diagram TDS vs N/(Na⁺+Ca²⁺) (Fig. 2a), the waters of Oyouloum and Manpal fall, as Lake Chad, within the rock-water interaction dominant zone. The waters of the other ouadis fall within the evaporation dominance zone. This indicates that water chemistry is primarily controlled by the evaporation-crystallization process. The evaporation process increases from the Moyolo ouadi to the Youlou ouadi, with the TDS being slightly lower than that of seawater. The Konorom, Barkadroussou, Touffou, and Yilingui ouadis have TDS values higher than that of seawater.

In the Gibbs diagram TDS vs Cl⁻/(Cl⁻+ HCO₃⁻) (Fig. 2b), the waters of Oyouloum, Manpal, and Lake Chad also fall within the rock-water interaction dominant zone. The water chemistry



of the other ouadis falls above the evaporation dominance zone defined by Gibbs (1970), which does not take into account soda lakes.

The Pearson correlation coefficient, r, shows a strong positive correlation (p < 0.05) of Na⁺ (r = 1), CO₃²⁻ (r = 0.97), HCO₃⁻ (r = 0.95), SiO₂ (r = 0.96), K⁺ (r = 0.92), PO₄²⁻ (r = 0.92), Mg²⁺ (r = 0.88), Cl⁻ (r = 0.84), and Br⁻ (r = 0.84) with TDS. The concentration of these ions depends on the water concentration, which increases with the progression of evaporation. The other major ions show no correlation with TDS. Therefore, their concentrations are not influenced by evaporation.

4.5 Type Waters

The chemical analysis results of major ions have been plotted on the diagram proposed by Piper (1944) and Chadha (1999) to identify water types. We have also included the average concentrations of analyzed waters available in the literature (Iltis 1971a, 1971b; Maglione 1969, 1976). The diagrams were made using Simler's (2023) software.

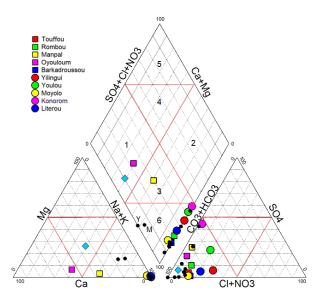


Fig.3. Piper diagram. Water types: 1, Ca-HCO3; 2, mixed Na-Cl-SO4; 3, mixed Ca-Na-HCO3; 4, mixed Ca-Mg-SO4-Cl, 5, Ca-Cl; 6, Na-HCO3. Blue diamond: Lake Chad. Analyze available in the literature: Y, Yoursoula; M, Moilo; Black solid circles, other ouadis (Iltis 1971a, 1971b; Maglione 1969; 1976).

The Piper diagram (Fig. 3) shows that the ouadi waters are predominantly of sodium bicarbonate type except for the freshwaters of Oyouloum, which is of calcium bicarbonate type, and Manpal, which is a mixed type of calcium bicarbonate and sodium bicarbonate.

The Chadha diagram divides water into eight types, as illustrated in Fig. 4. In this diagram, weak acidic anions $(HCO_3^-$ and CO_3^{2-} exceed strong acidic anions $(Cl^-$ and $SO_4^{2-})$ and, except for the Oyouloum ouadi, alkali metals $(Na^+ \text{ and } K^+)$ exceed alkaline earths $(Ca^{2+} \text{ and } Mg^{2+})$. The freshwater of the Oyouloum ouadi and Lake Chad belong to the Ca-Mg-HCO3 type. The other ouadis belong to the Na-HCO3 type.

These results are consistent with the older analyses available in the literature, which show that the ouadi waters are

predominantly of sodium bicarbonate type (Figs. 3 and 4). In the Piper and Chadha diagrams, the freshwater of Oyouloum, Manpal, Yoursoula (Maglione 1969), and Moilo (Maglione 1976) stand out from the other ouadis due to their lower sodium content. The low salinity of certain ouadis can be explained by leakage of salts into the underlying groundwater, as is the case for Lake Chad, and also by a significant salinity gradient from the edge of the ponds fed by freshwater sources to center, saltier due to evaporation (Maglione 1976).

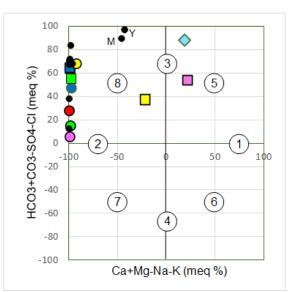


Fig. 4. Chadha diagram. Same legend as Fig.3. (1) Alkaline earths exceed alkali metals, (2) alkali metals exceed alkaline earths, (3) weak acidic anions exceed strong acidic anions, (4) strong acidic anions exceed weak acidic anions, (5) Ca–Mg–HCO3 water type, (6) Ca–Mg–Cl water type, (7) Na–Cl water type, (8) Na–HCO3 water type.

4.6 Trace Elements

The concentrations of trace elements are presented in Table 1. To assess the enrichment of trace elements during the evaporation of ouadi waters, we compare the trace element concentrations to those of freshwater from the Southern part of Lake Chad. The normalized trace element profiles of ouadi waters (Fig. 5) reveal diverse enrichments and depletions depending on the elements and the considered ouadis.

The lithium (Li) concentrations $(0.92-27.75 \ \mu g/L)$ are higher than that of Lake Chad (0.30 $\mu g/L$). Most ouadis have concentrations <10 $\mu g/L$, comparable to the soda Lake Turkana (Otachi et al., 2015). However, the ouadis of Barkadroussou, Touffou, and Yilingui have higher concentrations (16.54–20.05 $\mu g/L$), which are still much lower than those of Li-exploited salt lakes (Labbé and Daw, 2012).

Rubidium (Rb) concentrations in Oyouloum (1.83 μ g/L) and Manpal (1.65 μ g/L) are lower than that of Lake Chad (6.56 μ g/L) and comparable to the soda Lake Turkana (<10 μ g/L) (Otachi et al., 2015). Other ouadis exhibit more variable Rb concentrations (16.58–7,242 μ g/L), some of which are lower or within the range of salt lake concentrations (300–16,900 μ g/L) (Gao et al., 2020).

Cesium (Cs) concentrations in ouadis are either lower or equal $(0.01-0.02 \text{ }\mu\text{g/L})$ to that of Lake Chad $(0.02 \text{ }\mu\text{g/L})$ or



higher (0.06–1.37 μ g/L). These concentrations are low compared to salt lakes (34–2,500 μ g/L) (Gao et al., 2020).

Beryllium (Be) concentrations in ouadis ($<0.003-0.05 \mu g/L$) are very low.

Strontium (Sr) concentrations in Oyouloum (89.74 μ g/L) and Rombou (44.10 μ g/L) are lower than that of Lake Chad (91.23 μ g/L). Other ouadis show higher concentrations (97.97–395 μ g/L), comparable to soda lakes Turkana, Nakuru, and Bogoria (10–286 μ g/L) (Jirsa et al., 2013; Otachi et al., 2015).

Barium (Ba) concentrations in Oyouloum (32.76 μ g/L), Manpal (29.82 μ g/L), and Rombou (22.75 μ g/L) are lower than that of Lake Chad (70.38 μ g/L). Other ouadis have higher concentrations (157.70–255.80 μ g/L), which fall within the range of soda lakes Turkana, Nakuru, and Bogoria (<10–643 μ g/L) (Jirsa et al., 2013; Otachi et al., 2015).

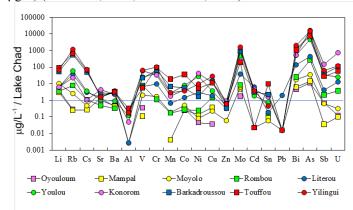


Fig. 5. Profiles of trace elements normalized with the values of the Southern part of Lake Chad.

Aluminum (Al) concentrations ($<0.129-15.04 \ \mu g/L$) are lower than that of Lake Chad ($47.87 \ \mu g/L$). These concentrations are low compared to the soda Lake Turkana ($860 \ \mu g/L$) (Otachi et al; 2015), the salt lakes of South Africa ($>150 \ \mu g/L$) (Melho et al., 2023), and the Altai region ($170-720 \ \mu g/L$) (Leonova et al., 2007).

Vanadium (V) concentrations in Oyouloum (1.33 μ g/L) and Manpal (0.2 μ g/L) are lower than that of Lake Chad (3.88 μ g/L). Other ouadis exhibit higher concentrations (7.83–245.90 μ g/L) compared to Lake Chad. These values are comparable to those of South African salt lakes (5–147 μ g/L) (Melho et al., 2023).

Chromium (Cr) concentrations in Oyouloum and Manpal are lower than the detection threshold of 0.012 μ g/L. Other ouadis show concentrations (1.69–100.80 μ g/L) higher than that of Lake Chad (1.00 μ g/L). These concentrations are comparable to those of the Altai salt Lakes (11-16 μ g/L) (Leonova et al., 2007) and the Urubu salt Lake (50 μ g/L) (Akubugwo et al., 2007).

Manganese (Mn) concentrations in Oyouloum are below the detection threshold (0.003 μ g/L). Ouadis of Manpal, Moyolo, Rombou, and Literou exhibit lower concentrations (0.01–2.31 μ g/L) than that of Lake Chad (3.43 μ g/L). Other ouadis show higher concentrations. These concentrations are within the range of soda lakes Lonar (0.07 μ g/L) (Surve et al., 2020) and Turkana (28 μ g/L) (Otachi et al., 2015).

Cobalt (Co) concentrations are diverse. Ouadis of Oyouloum, Manpal, Moyolo, and Rombou have concentrations $(0.07-0.14 \ \mu g/L)$ lower than that of Lake Chad $(0.28 \ \mu g/L)$. Other ouadis show concentrations $(0.41-9.98 \ \mu g/L)$ higher. These concentrations are low compared to the Sambhar soda lake $(2,450-9,100 \ \mu g/L)$ (Cherakar and Pathak, 2016; Pathak and Cherakar, 2015).

Nickel (Ni) concentrations are varied. Ouadis of Oyouloum, Manpal, Moyolo, and Rombou have concentrations ($0.08-0.45 \mu g/L$) lower than that of Lake Chad ($1.89 \mu g/L$). These low concentrations are comparable to those of the soda lake Turkana (<5 $\mu g/L$) (Otachi et al., 2015). Other ouadis exhibit concentrations ($5.73-78.61 \mu g/L$) higher than Lake Chad and within the range of South African salt lakes ($2-8,000 \mu g/L$) (Melho et al., 2023).

Copper (Cu) concentrations are varied. The ouadis of Oyouloum, Manpal, and Moyolo have concentrations (0.08–0.81 μ g/L) lower than that of Lake Chad (2.18 μ g/L) and comparable to those of the soda Lake Turkana (<0.2 μ g/L) (Otachi et al., 2015) and Lonar (0.45 μ g/L) (Surve et al., 2021). Other ouadis show higher concentrations (4.09–60.51 μ g/L), comparable to those of salt lakes in the Altai region (8–13 μ g/L) (Leonova et al., 2007) and the Urubu Salt Lake (15 μ g/L) (Akubugwo et al., 2007).

Zinc (Zn) concentrations (0.97–14.99 μ g/L) are lower than that of Lake Chad (16.41 μ g/L). These concentrations fall within the range of soda lakes Nakuru and Sambhar (<10–600 μ g/L) (Jirsa et al., 2013; Pathak and Cherakar, 2015).

Molybdenum (Mo) concentrations show diversity $(0.75-691.90 \ \mu g/L)$ and are higher than that of Lake Chad $(0.43 \ \mu g/L)$. These concentrations are within the range of soda lakes Van and Nakuru $(0.02-1,712 \ \mu g/L)$ (Reimer et al., 2009; Jirsa et al., 2013).

Silver (Ag) concentrations are below the detection limit $(0.009 \ \mu g/L)$.

Cadmium (Cd) concentrations in Oyouloum, Manpal, Moyolo, and Rombou ($<0.003-0.003 \ \mu g/L$) are lower than that of Lake Chad (0.13 $\mu g/L$). Other ouadis exhibit higher concentrations (0.25–1.65 $\mu g/L$). These low concentrations are similar to those of unpolluted lakes for Cd, such as the soda lake Turkana ($<5 \ \mu g/L$) (Otachi et al., 2015) or the Urubu salt lake (3 $\mu g/L$) (Akubugwo et al., 2007).

Ouadis of Konorom, Barkadroussou, and Touffou have tin (Sn) concentrations (7.8–47.38 μ g/L) higher than that of Lake Chad (5.00 μ g/L). Other ouadis exhibit lower concentrations (0.29–3.63 μ g/L) than that of Lake Chad.

Ouadis of Literou, Konorom, Barkadroussou, and Touffou have lead (Pb) concentrations ($0.38-6.13 \mu g/L$) higher than that of Lake Chad ($0.19 \mu g/L$). Other ouadis exhibit very low concentrations ($<0.003-0.003 \mu g/L$) of Pb.

Bismuth (Bi) concentrations $(25.59-11,400 \ \mu g/L)$ show significant variability and are higher than that of Lake Chad $(5.69 \ \mu g/L)$.

Arsenic (As) concentrations $(0.49-697.70 \ \mu g/L)$ show considerable variability and are higher than that of Lake Chad $(0.04 \ \mu g/L)$. Arsenic concentrations in Youlou, Konorom, Barkadroussou, Touffou, and Yilingui (201.30–697.70 $\mu g/L)$ are elevated compared to those of soda lakes of Sambhar (20



 μ g/L) (Pathak and Charaka, 2015), Nakuru (9–103 μ g/L), and Bogoria (18–132 μ g/L) (Jirsa et al., 2013). These high concentrations of As pose problems for spirulina harvested in the ouadis due to the ability of spirulina to absorb heavy metals as well as toxic elements such as As (Gershwin and Belay 2008; Doshi et al., 2009; Ahmad et al., 2010). Arsenic concentrations exceeding the tolerable limit of 1 ppm for human consumption (USP, 2015) have been reported in spirulina from Chad (Vicat et al., 2014).

Antimony (Sb) concentrations in Oyouloum, Manpal, and Moyolo ($0.002-0.04 \ \mu g/L$) are lower than that of Lake Chad ($0.04 \ \mu g/L$). Other ouadis show concentrations ($0.12-3.55 \ \mu g/L$) which are low compared to those of salt lakes in South Africa (5–200 $\mu g/L$) (Melho et al., 2023).

Uranium (U) concentrations in Oyouloum, Manpal, and Moyolo ($0.03-0.049 \ \mu g/L$) are lower than that of Lake Chad ($0.04 \ \mu g/L$). Other ouadis exhibit concentrations ($3.49-201.40 \ \mu g/L$) higher and comparable to those of the Sambhar salt lake ($8-1,400 \ \mu g/L$) (Yadav and Sarin, 2009).

The Pearson correlation coefficient, r, shows a strong positive correlation (p < 0.05) with the TDS for Bi (r = 1), Cu (r = 0.98), Rb (r = 0.94), Cr (r = 0.93), Cs (r = 0.89), Mo (r = 0.88), As (r = 0.84), Li (r = 0.79), and a good positive correlation for V (r = 0.74) and Al (r = 0.67). These elements tend to be conserved in the ouadi waters throughout the evapoconcentration process. The other trace elements do not show a correlation with the TDS.

V. CONCLUSION

The chemical composition of the waters of the Kanem ouadis is that of soda lakes with varying salinity and alkalinity levels. The concentrations of major ions show high variability compared to Lake Chad. Water chemistry is primarily influenced by the process of evapoconcentration. The predominant water type is sodium bicarbonate, consistent with historical analyses.

This study presents the initial data on trace element contents. Trace elements display varying degrees of enrichment or depletion when compared to Lake Chad. Trace element concentrations are similar to those found in soda lakes or saline lakes.

The concentrations of Na⁺, K⁺, Mg²⁺, HCO₃⁻, CO₃²⁻, Cl⁻, PO₄²⁻, Br⁻, SiO₂, Li, Rb, Cs, Cr, Cu, Mo, Bi, As, Al and V are positively correlated with the evapoconcentration process.

Due to sometimes high concentrations of As in water, this study recommends that spirulina development projects in the Kanem region include monitoring this toxic element in spirulina harvested in the ouadis.

ACKNOWLEDGMENTS

The authors thank the Faculty of Exact and Applied Sciences of the University of N'Djamena (Republic of Chad) for facilitating the field and laboratory work. They also thank the French Ministry of Foreign Affairs for funding the chemical analyses.

REFERENCES

- Ahmad A., Ghufran R., and Wahid Z.A. (2010). Metals Cd, As, Cu, and Zn Transfer through Dry to Rehydrated Biomass of Spirulina Platensis from Wastewater Polish J. of Environ. Stud., 19 (5): 887-893.
- [2] Akubugwo I.E., Ofoegbu C.J., and Ukwuoma C.U. (2007). Physicochemical Studies on Uburu Salt Lake Ebonyi State-Nigeria. Pakistan Journal of Biological Sciences, 10: 3170-3174.
- [3] Amaral P.G.C., Vincens A., Guiot J., Buchet G., Deschamps P., Doumnang J.C., and Sylvestre F. (2013). Palynological evidence for gradual vegetation and climate changes during the "African Humid Period" termination at 13°N from a Mega-Lake Chad sedimentary sequence. Climate of the Past 9:1–19. http://dx.doi.org/10.5194/cp-9-223-2013
- [4] Appelo, C.A.J. and Postma C. (1996). Geochemistry, groundwater, and pollution. A.A. Balkam ed., Rotterdam, Netherlands.
- [5] Chadha, D.K. (1999). A Proposed New Diagram for Geochemical Classification of Natural Waters and Interpretation of Chemical Data. Hydrogeol. J., (7): 431–439.
- [6] CBLT (2023). Note d'information sur la situation hydrologique dans le bassin du Lac Tchad, présenté au 2^{ème} Forum International Annuel sur le Développement de la Région du Lac Tchad, Commission du Bassin du Lac Tchad, Niamey, Niger, 2023.
- [7] Cherekar M.N. and Pathak A. P. (2016). Chemical assessment of Sambhar Soda Lake, a Ramsar site in India. Journal of Water Chemistry and Technology, 38(4):244-247.
- [8] Doumnang Mbaigane J.C. (2008). "Rapport de mission du 27 au 31 octobre 2007 : échantillonnage des sols, des eaux et des spirulines des ouadis du Kanem et du Lac", Geol. Lab., Univ of N'Djamena, Chad, Tech. Rep. TR-23/2008.
- [9] Doshi H.V., Ray A., and Kothari I.L. (2009). Live and dead Spirulina sp. to remove arsenic (V) from water. International Journal of Phytoremediation 11(1):53-64.
- [10] Dubrovsky, N.M., Burow, K.R., Clark, G.M., Gronberg, J.M., Hamilton P.A., Hitt, K.J., Mueller, D.K., Munn, M.D., Nolan, B.T., Puckett, L.J., Rupert, M.G., Short, T.M., Spahr, N.E., Sprague, L.A., and Wilber, W.G. (2010). The quality of our nation's waters— Nutrients in the nation's streams and groundwater, 1992–2004: U.S. Geological Survey, Reston, Virginia. Circular 1350, 174 p.
- [11] Dupont R. (1967). Etude des formations sédimentaires du Kanem. Premiers résultats. Office de la Recherche Scientifique et Technique Outre-Mer, Centre de Fort-Lamy, 150 p.
- [12] Dussart, B. (1966). Limnologie. L'étude des eaux continentales. Gauthier-Villars & Cie éd, Paris, 704 p.
- [13] Forrest, F., Rodvang, J., Reedyk, S. and White, J. (2006). A survey of nutrients and major ions in shallow groundwater of Alberta's agricultural areas. Prepared for the Prairie Farm Rehabilitation Administration Rural Water Program, Project Number:4590-4-20-4. Alberta Agriculture, Food and Rural Development, Edmonton, Alberta, 116 p.
- [14] Gao L., Ma G., Zheng Y., Tang Y., Xie G., Yu J., Liu B., and Duan J. (2020). Research Trends on Separation and Extraction of Rare Alkali Metal from Salt Lake Brine: Rubidium and Cesium, Solvent Extraction and Ion Exchange, 38,7:753-776. https://doi.org/10.1080/07366299.2020.1802820
- [15] Gershwin, M. E. and A. Belay A. (2008). Spirulina in human nutrition and health. CRC Press: 328 p.
- [16] Gibbs R.J. (1970). Mechanisms controlling world water chemistry, Science 170:1088-1090.
- [17] Hammer U.T. (1986). Saline Lake Ecosystems of the World. Monographiae Biologicae. Dr. W. Junk Publishers, Dordrecht, Boston, x, 616 p.
- [18] Iltis A. (1968). Tolerance de salinité de Spirulina Platensis (Gom.) Geitl, (Cyanophyta) dans les mares natronées du Kanem (Tchad). Cah. O.R.S.T.O.M, sér. Hydrobiol II (3-4):119–125.
- [19] Iltis A. (1969). Phytoplancton des eaux natronées du Kanem (Tchad).
 1, les lacs permanents à spirulines Cah. O.R.S.T.O.M, sér. Hydrobiol III (2):29–44.
- [20] Iltis A. (1971a). Phytoplancton des eaux natronées du Kanem (Tchad). V. Les lacs mésohalins. Cah. O.R.S.T.O.M., sér. Hydrobiol., V(1): 73-84.



- [21] Iltis A. (1971b). Note sur Oscillatoria (sous-genre Spirulina) Platensis (Nordst.) Bourrelly (Cyanophyta) au Tchad. Cah. O.R.S.T.O.M, sér. Hydrobiol, V(I): 53-7
- [22] Jirsaa F., Gruberb, M., Stojanovica A., Omondic, S.O., Maderd D., Körnere W., and Schagerlb M. (2013). Major and trace element geochemistry of Lake Bogoria and Lake Nakuru, Kenya, during extreme drought. Chemie der Erde, 73:275–282.
- [23] Kharaka Y. K. and Hanor J. S. (2003). Deep fluids in the continents: I. Sedimentary basins, (5): 499-540.
- [24] Kempe, S. and Kazmierczak, J. (2011). Soda Lakes. In: Reitner, J., Thiel, V. (eds) Encyclopedia of Geobiology, p. 824-828. Encyclopedia of Earth Sciences Series. Springer, Dordrecht. https://doi.org/10.1007/978-1-4020-9212-1_191
- [25] Labbé J.F. et Daw G. (2012). Panorama 2011 du marché du Lithium. Rapport BRGM RP/61340-FR, 151 p.
- [26] Leblanc M., Favreau G., Maley J., Nazoumou Y., Leduc C., Stagnitti F., van Oevelen P.J., Delclaux F., and Lemoalle J. (2006a). Reconstruction of Megalake Chad using Shuttle Radar Topographic Mission data, Palaeogeogr. Palaeoclim. Palaeoecol 239:16–27.
- [27] Leblanc M., Leduc C., Stagnitti F., Van Oevelen PJ., Jones C., Mofor L.A., Razack M., and Favreau G. (2006 b). Evidence for Megalake Chad, north-central Africa, during late Quaternary from satellite data, Palaeogeogr. Palaeoclim. Palaeoecol 230:230–242.
- [28] Leonova G.A., Bobrov V.A., Bogush A.A., Bychinskii V.A., and Bychinskii G.N. (2007). Geochemical Characteristics of the Modern State of Salt Lakes in Altai Krai. Geochemistry International, (45) 10: 1025–1039.
- [29] Maglione G. (1968). Présence de gaylussite et de trona dans les « natronières » du Kanem (pourtour nord-est du Lac Tchad). Bull Soc fr Minéral Cristallogr 9I:388–395.
- [30] Maglione G. (1969). Premières données sur le régime hydrogéochimique des lacs permanents du Kanem (Tchad). Cah. O.R.S.T.O.M., sér. Hydrobiol III (1):121–141.
- [31] Maglione G. (1971). Un exemple de comportement de la silice en milieu confiné carbonaté sodique : les natronières du Tchad. Bulletin du Service de la carte géologique d'Alsace et de Lorraine 24(4):255– 268.
- [32] Maglione G. (1976). Géochimie des évaporites et silicates néoformés en milieu continental confiné. Les dépressions interdunaires du Tchad – Afrique. Travaux et documents de l'O.R.S.T.O.M., n° 50, 340 p.
- [33] Maglione G. and Karn M. (1975). Spectres infrarouges des minéraux salins et des silicates néoformes dans le bassin tchadien. Cah. O.R.S.T.O.M., sér. Géol VII (1):3–9.
- [34] Maley J. (2010). Climate and Palaeoenvironment evolution in north tropical Africa from the end of the Tertiary to the Upper Quaternary. Palaeoecology of Africa 30:227–278.
- [35] Mehlo B., Lewis A., and Chivavava J. (2023). Comprehensive characterization of selected South African brines. Water SA, 49(2):126–135. https://doi.org/10.17159/wsa/2023.v49.i2.3957
- [36] OEHHA (1997). Public health goals for nitrate and nitrite in drinking water. Office of Environmental Health Hazard Assessment, California Environmental Protection Agency, Sacramento, California. Available at: www.oehha.ca.gov/water/phg/pdf/nit2_c.pdf.
- [37] Otachi E.O., Plessl C., Körner W., Avenant-Oldewage A., and Jirsa F. (2015). Trace Elements in Water, Sediments and the Elongate Tigerfish Hydrocynus forskahlii (Cuvier 1819) from Lake Turkana, Kenya Including a Comprehensive Health Risk Analysis. Bull Environ Contam Toxicol, 95(3):286-91.
- [38] Pathak A.P. and Cherekar M.N. 2015.Hydrobiology of hypersaline Sambhar Salt Lake a Ramsar site, Rajasthan, Indian Journal of Geo-Marine Sciences, 4(10) 640-1645.

- [39] Piper A. M. (1944). A graphic procedure in the geochemical interpretation of water-analyses. Eos, Transactions American Geophysical Union, 25(6): 914–928.
- [40] Reimer A., Landmann G., and Kempe S. (2009). Lake Van, Eastern Anatolia, Hydrochemistry and History. Aquat Geochem (2009) 15:195–222.
- [41] Sarath Prasanth S.V., Magesh N.S., Jitheshlal K.V., Chandrasekar N., and Gangadhar K. Evaluation of groundwater quality and its suitability for drinking and agricultural use in the coastal stretch of Alappuzha District, Kerala, India. (2012). Appl. Water Sci.,2: 165– 175.
- [42] Schneider J.L. (1967 a). Relations entre le lac Tchad et la nappe phréatique. Ex. Publ. nº 70 de l'A.I.H.S. Symposium de Garda, 122-131.
- [43] Schneider J.L. (1967 b). Carte hydrogéologique de la République du Tchad au 1/500 000e. Feuille de Mao. Rapport de synthèse. Rap. B.R.G.M., Paris.
- [44] Schuster M., Roquin C., Duringerr P., Brunet M., Caugy M., Fontugne M., Mackaye Taïsso H., Vignaud P., and Ghienne J.F. (2005). Holocene Lake Mega-Chad palaeoshorelines from space. Quaternary Science Review 24:1821–1827.
- [45] Servant M. (1983). Séquences continentales et variations climatiques : évolution du bassin du Tchad au Cénozoïque supérieur. Travaux et documents de l'ORSTOM, 159, 567 p.
- [46] Servant M. and Servant S. (1970). Les formations lacustres et les diatomées du Quaternaire récent du fond de la cuvette tchadienne. Rev. Geogr. Phys. Geol. Dy 12:63–75.
- [47] Simler R. (2023). Diagrammes (version 6.7) [Logiciel]. Obtenu à partir de http://www.lha.univ-avignon.fr
- [48] Surve R.R., A.V. Shirke A.V., Athalye R.R., and Sangare M.M. (2021). A Review on Chemical and Ecological Status of Lonar Lake. Current World Environment. 16 (1): 61-69.
- [49] Sylvestre F. (2014). Variabilité paléohydrologique et changements climatiques. In IRD (ed) Le développement du lac Tchad : situation actuelle et futurs possibles, Marseille, pp. 79–92.
- [50] Sylvestre F., Deschamps P., Sinine A.B., Rirongarti R., Mazur J.C., Waldmann N., Paula Do Amaral P., Bouchez C., and Doumnang J.C. (2019). Paléoenvironnements et variations paléohydrologiques du lac Tchad au cours des 12 000 dernières années. In : IRD (ed) Le Tchad des lacs. Les zones humides sahéliennes au défi du changement global, chap 2, pp. 53–64.
- [51] Touchard L. (2000). Qu'est-ce qu'un Lac ? (What is a lake ?). Bull Assoc Géogr Franç, 77(4):313-322.
- [52] UNESCO (1994). The major ion composition of seawater. US Department of Energy, version 2, September 1994, ORNL/CDIAC-74, chap. 5, part 6.1, pp. V-10.
- [53] USP (2015). United States pharmacopeia dietary supplements. United States Pharmacopeial Convention edit, 28 juil. 2015, 3795 p.
- [54] Vicat J.P., Doumnang Mbaigane J.C., Bellion Y. Teneurs en éléments majeurs et traces de spirulines (*Arthrospira platensis*) originaires de France, du Tchad, du Togo, du Niger, du Mali, du Burkina-Faso et de République centrafricaine. C.R. Biologie, 337: 44–52.
- [55] WHO (2017). Guidelines for Drinking-Water Quality, 4th ed. Incorporating the First Addendum. World Health Organization, Geneva, Switzerland, 541 p.
- [56] Yadav D.N. and Sarin M.M. (2009). Geo-chemical behavior of uranium in the Sambhar Salt Lake, Rajasthan (India): Implications to "Source" of salt and uranium "Sink". Aquat. Geochem, (15):529– 545.